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Utilisation of CryoSat-2 SAR altimeter in operational ice charting

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CryoSat-2 operational ice charting

E. Rinne and M. Similä

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Abstract

We present methods to utilise Cryosat-2 (CS-2) Synthetic Aperture (SAR) mode data in operational ice charting. We compare CS-2 data qualitatively to Synthetic Aperture Radar (SAR) mosaics over Barents and Kara seas. Furthermore, we compare the CS-2 to archived operational ice charts. We present distributions of four CS-2 waveform parameters for different ice types as presented in the ice charts. We go on to present an automatic classification method for CS-2 data which, after training with operational ice charts, is capable of determining open water from ice with a hit rate of > 90%. The training data is dynamically updated every five days using the most recent 15 days CS-2 data and operative ice charts. This helps the adaption of the classifier to the evolving ice/snow conditions throughout winter. The classifier is also capable of detecting three different ice classes (thin and thick first year ice as well as old ice) with success rates good enough for the output to be usable to support operational ice charting. Finally, we present a near real time CS-2 product just plotting the waveform characteristics and conclude that even such a simple product is usable for some of the needs of ice charting.

1 Introduction

Our aim is to present new methods to utilise satellite altimeter measurements in operational ice charting. We present an automatic classification method to derive different ice stages of development from Cryosat-2 (CS-2) waveforms. This is different from the most common sea ice application of satellite altimeters today, which is measuring the freeboard and the thickness of Arctic winter sea ice, see for example (Laxon et al., 2013).

The use of altimeters to support ice mapping has been suggested already 35 years ago by Dwyer and Godin (1980). Ice detected utilising altimeter waveforms has been compared with sea ice extents from passive microwave satellite instruments (Laxon,

TCD

9, 4117–4145, 2015

Cryosat-2 operational ice charting

E. Rinne and M. Similä

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Cryosat-2 operational ice charting

E. Rinne and M. Similä

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



1990) and ERS-1 altimeter based sea ice estimates were faxed to research vessels navigating in the Southern seas in the early 90's (Laxon, 1994). Mostly due to Synthetic Aperture Radars (SAR) becoming the standard tool in operational sea ice charting, altimetry has developed into a method for climate research (Laxon et al., 2003, 2013; Giles et al., 2008; Kwok et al., 2009). Altimetry is, however, widely used in numerical weather prediction where fast delivery products from different altimeters on open ocean are assimilated into weather models (Vidard et al., 2009).

The diminishing Arctic sea ice cover is opening new sea routes. In consequence, navigation in seasonally ice covered waters is due to increase rapidly. This calls for accurate and timely sea ice information, especially on dangerous sea ice conditions. Vessels navigating in or near sea ice can be roughly divided into those wanting to completely avoid ice and those that can safely operate in medium first year ice. For the ice avoiding ships, ice edge detection is enough but for the latter group some information on the stage of development of the ice is needed as well. Most widely used Earth Observation (EO) data in operational ice charting are Synthetic Aperture Radar (SAR) frames. In areas of heavy traffic, such as the Baltic Sea for example, SAR data is virtually indispensable. However, because the number of SAR acquisitions is limited, other instruments, such as altimeters, may provide valuable additional information on sea ice.

Due to the number of SAR satellites flying today being reasonably small, ice services can face incidents when fresh SAR data is simply not available. The loss of Envisat and its ASAR instrument in April 2012 pointed out how dependent European ice charting was on a single satellite. Furthermore, SAR images have a limited spatial coverage which results in data gaps. For ice covered seas with little or only sporadic traffic, SAR frames are often not acquired. This is the case with Southern ice covered seas – as the number of ground stations in the Antarctic is small, SAR acquisitions over Southern sea ice would use costly satellite mass storage. This is different in the Arctic seas: for example the SAR frames from Kara and Barents seas could be directly downlinked to a ground station in northern Europe. Thus it is convenient to study altimeter ice charting

in the European sector of the Arctic even if one of our aims is to contribute to Antarctic sea ice charting.

In absence of SAR data, ice services look for auxiliary data. Optical satellite images, such as MODIS or Suomi-NPP images may mitigate the problem, but only if the lighting and cloud conditions are favourable. Alas the polar night and frequent cloud cover often render optical imagery useless. In these cases ice services are bound to either use low resolution (at the best in 10 km grid) products from, for example the EUMETSAT OSI-SAF project (<http://saf.met.no/p/ice/>) or, in the worst case, inform the users that up to date ice information is not available.

Zygmuntowska (2014) examined the possibility to distinguish between first year and multi year ice throughout the winter using altimeter waveform parameters. They gridded their data in a 25 km grid and studied the parameters one at a time and found that the spatial distribution of a single parameter varied strongly from month to month from the freezing-up period to the Arctic spring. Sometimes the patterns followed the MY/FY ice areas, other times not. This we can regard as manifestation of sensitivity of the altimeter waveform for the ice/snow surface conditions.

2 Data

Our study area is the Barents and Kara seas in the European sector of the Arctic Mediterranean. We chose the area because of the good availability of data, namely the Arctic and Antarctic Research Institute (AARI) operational ice charts and an archive of SAR data as a heritage from the Enhanced Arctic Sea Ice Information – ANISTIAMO exercise carried out by FMI in 2014 (ANISTIAMO Reports, 2014) and its predecessor described in Similä et al. (2013).

Cryosat-2 operational ice charting

E. Rinne and M. Similä

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



2.1 Cryosat-2 data

We use the CS-2 SAR mode (Wingham et al., 2006) Level 1b (L1b) data (Bouzinac, 2014) available online from ESA. For this study we use the Near Real Time CS-2 products that were made available by ESA for our study.

5 The main instrument of CS-2 is the Synthetic aperture radar/Interferometric Radar Altimeter (SIRAL) altimeter. The CS-2 L1b product used here is, essentially, the CS-2 SAR-processed average waveform for each point along the ground track of the satellite. In the SAR mode, SIRAL employs the along-track beam formation (i.e. SAR processing) to generate a resolution cell of approximately 300 m by 1 km (Wingham et al., 2006). The SAR mode of CS-2 is originally designed for ice covered seas. With the along-track resolution enhancement due to SAR processing enables smaller leads to be detected within the sea ice pack as was possible with CS-2's predecessor, the Envisat RA-2 (Laxon et al., 2013). Thus the CS-2 is operated in SAR mode over the sea ice covered seas. This combined with the polar orbit of CS-2 with an inclination of 10 92° results in a good coverage of CS-2 SAR mode measurements over the ice covered Arctic oceans. 15

2.2 Ice charts

We take the ice stage of development from the weekly ice charts published by the Arctic and Antarctic Research Institute (AARI) (Bushuev and Loshchilov, 2007). The AARI ice charts are available online from <http://www.aari.ru/>. We downloaded the charts from the current AARI website predecessor, an AARI FTP-server, in SIGRID3 format (SIGRID3 Manual, 2014). The ice charts for the Kara and Barents Seas provide estimates of total ice concentration (CT), the partial concentrations (CA, CB and CC), and the stage of ice development (SA, SB and SC) for the three thickest ice types for polygonal areas. To 20 quantitatively study the effect of sea ice stage of development on altimeter waveforms, we rasterised the AARI ice maps into 2 km grids in Lambert equal area projection. We sampled these grids so that for each CS-2 measurement falling within the test 25

Cryosat-2 operational ice charting

E. Rinne and M. Similä

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



area the ice characteristics for the ice chart polygons were fetched from the temporally closest AARI map. As the AARI maps are generated weekly, the largest time difference between the the different data sets is three days.

The ice stage of development in the AARI ice charts follows the World Meteorological Organisation (WMO) sea ice nomenclature. The WMO nomenclature defines seven different stages of development: nilas, grey ice, gray-white ice, thin first year (FY) ice, medium FY ice, thick FY ice and old ice. Three first stages are thin ice, which are defined as ice that does not form pressure ridges. The three FY ice categories are ice which is thick enough to ridge and has not experienced a whole melt season. The last stage, old ice, is ice that has survived at least one whole summer melt season. In the cryospheric community, old ice is often referred to as multi year (MY) ice. In this study we use only three different stages of development: thin (< 70 cm) FY (WMO categories nilas, gray, gray-white and thin FY), thick (> 70 cm) FY (WMO categories medium and thick FY) and MY ice (WMO category old ice).

The AARI ice charts are based on SAR and optical satellite images as well as reports from coastal stations and ships. The segmentation of images and subsequent interpretation and mapping of ice conditions are carried out by ice experts. The main purpose of the weekly ice chart is to show the spatial distribution and characteristics of sea ice. We have not found English language scientific publications discussing the accuracy of the AARI weekly ice charts. However, they are the only source of ice information for this area with adequate spatial coverage and temporal resolution for our study. Furthermore they are completely independent of CS-2 data, which makes comparing the two meaningful.

The area of the polygons in the AARI ice chart varies largely. In some cases the smallest diameter of the polygon is around 10–20 km, in some cases the largest diameter of a polygon is up to 100–200 km. Even if the AARI charts provides partial concentrations for up to three ice types for each polygon, we assign just one ice type to a single polygon. In the Sect. 3.1. we need to apply threshold for the partial concentration of the dominant ice class. For the automatic classifier discussed in

Cryosat-2 operational ice charting

E. Rinne and M. Similä

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Sects. 3.1 and 4.2, this threshold is 75%. When we assign an ice class to a polygon in the test phase, no threshold is used and the partial concentration of the dominant ice type is often below 75%. Thus a large fraction of the polygon represents in reality some other ice development stage than what we have labelled it to be. In consequence, an inherent inaccuracy is present in our reference data.

2.3 SAR composites

To visualise CS-2 waveform characteristics, we use synthetic aperture radar composites compiled from Radarsat-2 frames. These composites were originally compiled to be used as input data for a multisensor ice thickness chart for the Barents and Kara seas in the ANISTIAMO demonstration project in spring 2014, see www.arcice.org. The SAR data in Figs. 1 and 2 is used for visualisation only and is included because it is the most widely used EO data in operational ice charting and we assume most of our readers to be familiar with it.

3 Methods

We chose four characteristics, pulse peakiness (PP), leading edge width (LEW), tail to peak power ratio (TPP) and stack standard deviation (SSD) to describe the CS-2 waveform. The PP, TPP and LEW are easily derived from the waveform and the SSD is delivered in the L1B data product (Bouzinac, 2014).

The PP is defined as:

$$PP = 128 \cdot \frac{\sum_{i=1}^{128} P_i}{P_{\max}} \quad (1)$$

where P_i is the power in the i th range bin and P_{\max} is the maximum power in one range bin in the waveform.



PP has been used previously to distinguish leads from ice floes in pulse limited radar altimeter data (Laxon et al., 2003). Classifying waveforms with high PP as leads is an integral step in traditional radar altimeter sea ice freeboard processing.

The SSD is taken from the CS-2 L1b data product. The SSD is the standard deviation of a set of Doppler echoes from a common surface location prior to stacking. SSD has been used, in conjunction with PP, in the lead detection, e.g. by Laxon et al. (2013); Ricker et al. (2014) and Kurtz et al. (2014).

For the LEW, we use the difference between the bins retracked with $\rho = 10\%$ and $\rho = 90\%$ percentage. Offset Center of Gravity (OCOG) retracker. The OCOG retracker returns the bin number where the received power count rises for the first time over the threshold value τ of:

$$\tau(\rho) = \frac{\rho}{100} \cdot \sqrt{\frac{\sum_{i=1}^{128} P_i^4}{\sum_{i=1}^{128} P_i^2}} \quad (2)$$

There the ρ is the percentage.

The OCOG retracker have been widely used in altimetry, see for example Wingham et al. (1986), Bamber (1994) or Soussi and Femenias (2006).

The TPP is defined as:

$$TPP = \frac{\frac{1}{21} \sum_{i=\max+50}^{\max+70} P_i}{P_{\max}} \quad (3)$$

where max is the index of the range bin with the maximum power.

TPP tells us how much off-nadir power there is present in the tail of the waveform. Typically TPP is high for surface with large roughness. We also studied the feasibility of another waveform characteristic suggested by Kurtz et al. (2014), but found that TPP yield better results in the automatic classification. The rationale of choosing is discussed in Sect. 4.2.

Cryosat-2 operational ice charting

E. Rinne and M. Similä

Title Page	
Abstract	Introduction
Conclusions	References
Tables	Figures
◀	▶
◀	▶
Back	Close
Full Screen / Esc	
Printer-friendly Version	
Interactive Discussion	



It is important to note here that we rely only on the four waveform characteristics: PP, SSD, TPP and LEW. We do not do freeboard processing in the style of Laxon et al. (2013) or use the freeboard values in the higher level CS-2 data products. This is because we want the CS-2 methods to be simple as well as independent from any other data sources.

3.1 Automatic classifier

We set out to build a classifier to retrieve the ice stage of development using only CS-2 waveforms and past AARI ice charts as input. Due to the reasons discussed below we selected a k -nearest neighbours (k -NN) classifier for our classification method. The adopted classifier is able to classify four different ice classes (open water, thin FY, thick FY and MY ice) with a reasonable accuracy.

The results of Zygmontowska (2014) lead us to adopt an approach where we dynamically update the training set for the classifier. The training data was gathered from the CS-2 acquisitions and the AARI charts during a 15 days period. Then we used these class boundaries in analysing the CS-2 data of the following five days period, called here test data. The training data and, hence, also the class boundaries, were recalculated at intervals of five days instead of applying fixed class boundaries all the time. The constantly updated class boundaries help the classifier to adapt to ice/snow conditions which evolve continuously.

As one can see from Figs. 3 and 4 the waveform parameters of different ice categories are often very close each other. In our data set this is true especially in November during the freeze-up period. It is obvious that we cannot separate the ice types using just one parameter. Instead a set of parameters should be utilised simultaneously. The shape of distribution of a single parameter for ice class deviated in most cases essentially from the normal distribution. The task to non-parametrically model a non-Gaussian multidimensional distribution is very challenging. Hence, we chose not to use a Bayesian classifier which is the theoretically optimal classification method (Hastie et al., 2001).

Cryosat-2 operational ice charting

E. Rinne and M. Similä

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Cryosat-2 operational ice charting

E. Rinne and M. Similä

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



We avoided the difficulty to deal with non-Gaussian multidimensional distributions by applying the k -nearest neighbors classifier. It's largest advantage is that it does not require a statistical model for the data. The classifier is memory based. The concept of the classifier is simple. We first collect all our training data together. When the class of a new feature vector from a test set is decided, we first search k closest training points and perform a majority vote among them. Ties are broken at random. The crucial requirements for the efficiency of the classifier are:

1. the training set must represent well the data to be classified,
2. dimension of the feature space can not be very large,
3. the number k of the neighbors must be determined soundly and
4. the distance between the measurements must be a proper metric.

We first discuss the distance metric. We scaled the distribution of each parameter on the interval $[0,2]$. The values of rare events were truncated. Naturally, we could have used also the unit interval $[0,1]$. Now the distance between feature vectors can be computed using the ordinary Euclidean distance.

During the 15 days training period we collect a rather extensive data set. In most cases this set covers the variation occurring in the following five days CS-2 data. To characterize the waveform we selected first the standard features (PP, LEW, SSD) as a part of a feature vector. Increasing the dimension of the feature space always reduces the amount of training points inside a unit cube. In consequence, a situation where a new feature vector does not have any training point close to it begins to occur more frequently and the classification accuracy decreases.

On the other hand, the discrimination between feature vectors is, in principle, more efficient if the description of a waveform is more versatile, i.e. the number of features is larger. Keeping in mind these these two conflicting consequences of adding features, we studied if an additional feature would indeed improve the results. We chose to

experiment with the two features suggested by Kurtz et al. (2014). The first one is the ratio of tail to peak power (denoted here Kf1):

$$Kf1 = \frac{\frac{1}{6} \sum_{i=\max+1}^{\max+6} P_i}{P_{\max}}, \quad (4)$$

This feature describes how rapidly the power P_{\max} declines immediately after the maximum value. The other feature proposed was the ratio of waveform tail to peak power or TPP as defined in Eq. (3) in Sect. 3.

A way to measure dependence between two random variables is to utilize the mutual information (MI) (Cover and Thomas, 1991). MI measures the overall dependency between two random variables, not just the linear dependency like the correlation coefficient. We calculated the pairwise values of MI between features Kf1 and TPP and the set of standard parameters (PP, LEW, and SSD). The MI values were consistently lower for TPP than Kf1. This implies that the TPP contains less overlapping information with the standard three parameters than Kf1.

Next we checked if the addition of TPP to the three previous parameters (PP, LEW and SSD) actually improves the k -NN classification. The result was that the addition of TPP increased the correct classification accuracy from 0 to 5% for a single class depending on the ice class and the test set. This was more than the improvement after addition of Kf1. We also tested a 5-dimensional (5-D) feature vector containing the parameters (Kf1, TPP, PP, LEW, SSD). The addition of Kf1 had a negative impact on the accuracy compared to the tested 4-dimensional (4-D) feature vector. Hence, we selected the features (PP, LEW, SSD, TPP) as our feature vector.

Prior to the classification we preprocess the data. All waveforms with LEW larger than 14 are excluded from analysis. We regard these waveforms too noisy to be useful. We also want to remove potential leads from data to limit the confusion between different ice types.

On the detection of potential leads we employ three slightly different statistics to characterize specular reflections. These three features consist of PP and two statistics

Cryosat-2 operational ice charting

E. Rinne and M. Similä

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



presented in Ricker et al. (2014). They are defined as follows:

$$PP_{\text{left}} = 9 \cdot \frac{P_{\text{max}}}{\sum_{i=\text{max}-1}^{\text{max}-3} P_i} \quad (5)$$

and

$$PP_{\text{right}} = 9 \cdot \frac{P_{\text{max}}}{\sum_{i=\text{max}+1}^{\text{max}+3} P_i} \quad (6)$$

where max refers to the index of the range bin with maximum power. We determined experimentally the thresholds: $PP > 40$ and either $PP_{\text{left}} > 20$ or $PP_{\text{right}} > 15$. Ricker et al. (2014) also used a SSD threshold for lead detection. However, in our data set when the PP conditions were met, SSD was almost always below 4 and we did not set an additional condition. The amount of potential lead signals was usually 5–7% from the measurements. Waveforms considered potential leads were excluded from the analysis.

To reduce speckle, we average the CS-2 waveform parameters over the five consecutive footprints and assign the value to a single footprint area, i.e. we use the running mean method for each parameter separately. All CS-2 measurements are subject to speckle like any other SAR signature. The speckle influence is most obvious for PP but it is also present in other waveform parameters. In this step we implicitly assume that the five consecutive footprints covering a track of 1900 m belong to the same ice class. Considering our coarse typing of ice classes (3–4 ice types) this is a reasonable assumption.

We classify the running means of the waveform parameters using the k -NN classifier. At the first phase we perform the classification for each 4-D feature vector separately. Then we take the mode of 50 consecutive ice class labels. The resulting mode is then regarded as the estimated ice type and it has a resolution of 19 km. If we take into account the spatial averaging the true resolution is about 20 km along the track. We

Cryosat-2 operational ice charting

E. Rinne and M. Similä

Title Page	
Abstract	Introduction
Conclusions	References
Tables	Figures
⏪	⏩
◀	▶
Back	Close
Full Screen / Esc	
Printer-friendly Version	
Interactive Discussion	



Cryosat-2 operational ice charting

E. Rinne and M. Similä

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



examined also the possibility to use only 30 consecutive class labels to achieve a better spatial resolution. In our test runs the 20 km resolution yielded, however, slightly more accurate results than the 12 km resolution. Although the difference was not large we chose the coarser resolution data because of slightly better reliability.

The mode operation is also applied for our ground truth data points extracted from the AARI ice charts (50 consecutive points corresponding the CS-2 points). When we assess the classification accuracy using the ground truth data, it takes place comparing these 19–20 km long segment tracks. In the classification maps in the Sect. 4.2 we show the results using a sliding window technique, i.e. the consecutive the class labels have just a distance of 0.38 km between them.

We still must determine a reasonable value for k . The value $k = 1$ yielded highly variable results for the test sets and was deemed unpractical for our purposes. When we compared the values $k = 3$ and $k = 5$ we noticed that the results were close to each other for the November test sets (3 sets) although $k = 3$ yielded slightly better results. In March the results more clearly favored $k = 3$ over $k = 5$. Increasing k to a larger value than $k = 5$ lead to poorer results. The difference between the November and March data sets was that the number of ice types was four in March and three in November. Hence the classification task in March was more challenging than in November. We have used the value $k = 3$ in our classifications.

The disadvantage of the k -NN classifier with the mode operation is that the classes with most measurements are usually best detected.

4 Results and discussion

4.1 Qualitative comparison

To visualise the behaviour of PP, Fig. 1 below shows PP drawn over a SAR composite from the same area on the 05 March 2014. The figure shows an increase from the low (PP < 3, blue) values over open ocean to high (PP > 7, red) values over ice. The

Cryosat-2 operational ice charting

E. Rinne and M. Similä

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



increase coincides with the ice edge visible in the SAR frame. Analogously to the Fig. 1 and PP, Fig. 2 shows the CS-2 SSD over a SAR frame. Again the ice edge is clearly visible as SSD changes from high values (SSD > 20, blue) over open water to small (SSD < 10, red) on ice. However, SSD seems to exhibit more variation over ice than PP does. There are areas of high SSD within the ice pack, some of which coincide with features in SAR data. The interpretation of a SAR frame to ice characteristics is subjective due to the multitude of factors affecting the backscattering. Areas of high backscatter can be thick heavily deformed ice or, in some cases, broken thin ice with varying amount of open water e.g. brash ice. Detailed information on SAR backscattering statistics in our test region can be found in Lundhaug (2002). Due to the ambiguity in SAR signature, it is impossible to derive the ice thickness from a single SAR frame alone. Because of this we do did not carry out a quantitative comparison of SAR and CS-2 data.

To further study the effect of ice stage of development on the CS-2 waveform, we sampled the AARI ice charts at a point closest to the CS-2 measurement as described in Sect. 2.2. The distributions of waveform parameters for different stages of development are shown in Figs. 3 and 4. Open water, as expected from the SAR/CS-2 comparison above, shows peaks in high (SSD > 50) SSD and low (PP < 5) PP. In March thin FY ice has a bimodal PP distribution, possibly due to the polygons labeled with thin ice often having a total ice concentration of less than 100%, ie. there are areas of open water present in addition to the ice. Overall, the distributions show promise for distinguishing different ice classes based on the four waveform parameters. The results from our automatic classifier are presented in the next subsection.

4.2 Automatic classification

The classification methodology is presented in Sect. 3.1. The waveform parameters are correlated but have also differences in their distributions as shown in Figs. 3 and 4. We utilised three different ice categories in November (open water, thin FY, and MY

ice) during the freeze-up period and four ice categories in the middle of winter in March (open water, thin FY, thick FY, and MY ice).

In the training data the dominant ice type from AARI charts was used as the true ice type for all CS-2 waveforms falling within the ice chart polygon if the partial concentration of the dominant ice type was > 75%. Only measurements from this kind of polygons were accepted for training. Then the CS-2 data for the five days test set (see Sect. 3.1) was classified using the system. The results of the CS-2 classification were then compared to the stages of development taken from the temporally closest AARI chart during the test period. In the test set we do not use the 75% rule for the dominant ice class, but use the stage of development which has the highest concentration as the truth for all of the CS-2 measurements falling within the polygon. Tables 1 and 2 show the classification matrices for November and March, respectively. Maps of the classification results are presented in Figs. 5 and 6 for November and March respectively.

In November thin FY ice, MY ice and open water are present both in the CS-2 data as well as in the AARI chart (Fig. 5 and Table 1). Thick FY ice was absent in the AARI charts we used as training for November, and thus we only have three classes: open water, thin FY and MY ice. The open water is classified right in 98% of the cases. The thin FY ice mixes somewhat with MY ice: 46% of CS-2 measurements in polygons marked to consist mostly of thin FY ice in the AARI charts are classified to be MY ice based on CS-2. Analogously 8% of CS-2 measurements from polygons where MY ice is dominant is classified as thin FY ice. Part of the inconsistency is natural. In reality there are inclusions of FY ice within the old ice area as well as inclusions of MY ice in the FY ice area. However, there are CS-2 measurements classified as MY ice south of 80° N. It is unlikely that these are in reality MY ice. We assume these to be areas of deformed ice where the large scale surface roughness is more akin to MY ice than recently formed FY ice. If this is the case, the information about deformed ice, most likely an obstacle to navigation, would be valuable for operational ice charting. Sadly, we have no means to test our assumption.

Cryosat-2 operational ice charting

E. Rinne and M. Similä

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Cryosat-2 operational
ice charting**

E. Rinne and M. Similä

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



For March (Fig. 6 and Table 2) the results are similar to November. The overall correspondence of AARI maps and the CS-2 classification is good. The two FY ice classes mix considerably. This is not an unsurprising result because the thickness of 70 cm is not a threshold which would abruptly change the characteristics of ice. Instead the division of ice thickness of less or larger than 70 cm is mostly based on the needs of ice navigation. Furthermore there are inclusions of thin FY ice within the thick FY and vice versa. The results for open water (93% right) and MY ice (83% right) are good. A notable feature in the chart is MY ice appearing in the CS-2 measurements at about 78° N and 95° E, West of Vilkitsy Strait. There were small amounts of MY ice near the coast marked in the AARI charts too, especially near the coasts. Thus it may be that our classification exaggerates the amount of MY ice, especially in the areas where heavy deformation is likely to occur. However, for the purposes of operational ice charting, a cautious approach is often preferred.

Our results are similar to those of Zygmontowska et al. (2013), obtained for an airborne altimeter in a smaller scale. Zygmontowska et al. (2013) used a Bayesian classifier and presented a comparison of CS-2 derived sea ice types and OSI-SAF ice types. They showed that ice type classification with satellite altimeter data is possible but also found regions where clear discrepancies occur between the CS-2 derived ice type and their validation data. They attributed these discrepancies to areas of FY ice with large surface roughness. We also found false positive MY ice classifications. As discussed before, for our application, classifying heavily deformed FY ice as MY ice is not a problem since both present similar threat to navigation.

When we inspect the classification results, the detection of thin ice (here FY ice thinner than 70 cm) has been least successful. As seen in Figs. 3 and 4 the waveforms originating from thin ice exhibit a wide range of variation for all used features. This is understandable. Very thin ice (thickness less than 10 cm) has often ice concentration well below 100%. Due to this, part of the waveforms assigned to thin ice are actually waveforms from open water. This can be seen especially clearly in the behaviour of SSD in Fig. 2. Processes such as rafting and ridging increase surface roughness, and

thin ice as defined in this work can be easily ridged. The thin ice is mostly mixed with thick FY ice (thickness larger than 70 cm) but to a lesser degree also with open water.

5 Cryosat-2 product to support FMI operational ice charting

We built a CS-2 ice product based on near real time waveform characteristics to test if the Finnish ice service would benefit from CS-2 data. The product is basically maps of Near Real Time (NRT) CS-2 PP, LEW and SSD. A comparably simple system was built to download the data from an ESA server and to calculate and plot PP, LEW and SSD. All of the CS-2 data acquired during the past two days is used for each product. The prototype system has been running at the Sodankylä satellite receiving station since September 2014. The FMI CS-2 product is available online at <http://ice.fmi.fi/Cryosat/>. The ice experts were told to look for abrupt changes in the PP to detect the ice edge and to interpret areas of constant low PP as open water. The analysts were also briefed that in addition to open water, heavily deformed sea ice may result in low PP. We did not endeavour to make an automated ice edge detection since the ground tracks are rather sparse and we did not want to interpolate between them.

The product was received well. During the fall the Finnish ice service provided ice information to ships in the Kara Sea and the CS-2 product was used as an independent reality check for Sentinel-1 and COSMO-Skymed SAR frames. As expected, in an area where SAR data is readily available, such as the Kara Sea, altimeter products add little or no value to operational ice charting. However, several requests were made to build a similar product for the Antarctic ice covered oceans. Our plan is to continue providing the FMI CS-2 product and implement a similar product for the Sentinel-3A data in the future.

In April 2015, University College of London (UCL) published a near real time CS-2 sea ice thickness product available online. The UCL product is based on the CS-2 processing chain presented in Laxon et al. (2013). At the moment of writing, the UCL

Cryosat-2 operational ice charting

E. Rinne and M. Similä

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



product lacks validation or comparison to independent EO and it is hard to assess its usability for navigation.

6 Conclusions

For the first time, we have demonstrated the use of a SAR altimeter, namely the SIRAL-2 onboard CS-2, to support operational ice charting. We've presented a qualitative comparison of a SAR composite and CS-2 data. Furthermore, we've compared the CS-2 waveform characteristics to the stage of development of ice taken from ice charts. We've also presented an automatic classification system capable of detecting open water, thin FY ice, thick FY ice and MY ice based on four CS-2 waveform characteristics. This is the first time ice classification methodology using satellite SAR altimeter data is presented and tested outside gray literature. The classification system requires recent operational ice charts for the training, but after the training the only input are CS-2 data. The system resolves MY ice and open water well. The two tested FY classes mix significantly, but for the application of operational ice charting this is not a problem.

We have built a prototype system providing simple maps of NRT CS-2 waveform characteristics. This product has been tested by the Finnish ice service during the winter of 2014–2015. The feedback was positive. Thus we conclude that satellite altimeters, in this case the CS-2, provide an independent source of sea ice information to complement SAR and passive microwave data. Albeit low resolution and sparse, altimeter measurements can be used at times and locations where other data sources are unavailable.

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Cryosat-2 operational ice charting

E. Rinne and M. Similä

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



References

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CryoSat-2 operational ice charting

E. Rinne and M. Similä

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Cryosat-2 operational ice charting

E. Rinne and M. Similä

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



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Discussion Paper | Discussion Paper | Discussion Paper | Discussion Paper | Discussion Paper

TCD

9, 4117–4145, 2015

Cryosat-2 operational ice charting

E. Rinne and M. Similä

Title Page	
Abstract	Introduction
Conclusions	References
Tables	Figures
◀	▶
◀	▶
Back	Close
Full Screen / Esc	
Printer-friendly Version	
Interactive Discussion	



Cryosat-2 operational ice charting

E. Rinne and M. Similä

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Table 1. Classification Matrix of CS-2 Based Classification and AARI Ice Chart, November 2013. Last column shows the best and the worst hit rates of the three 5-day periods.

	FY < 70 cm	MY	Open Water	Worst–Best
FY < 70	51 %	46 %	3 %	31–64 %
MY	8 %	92 %	0 %	88–97 %
Open Water	5 %	4 %	92 %	87–95 %

TCD

9, 4117–4145, 2015

Cryosat-2 operational
ice charting

E. Rinne and M. Similä

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

I◀

▶I

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Table 2. Classification Matrix of CS-2 Based Classification and AARI Ice Chart, March 2014. Last column shows the best and the worst hit rates of the three 5-day periods.

	FY < 70 cm	FY > 70 cm	MY	Open Water	Worst–Best
FY < 70	20%	60%	7%	13%	15–26%
FY > 70	1%	84%	14%	0%	75–91%
MY	0%	14%	86%	0%	77–92%
Open Water	2%	1%	3%	93%	90–98%

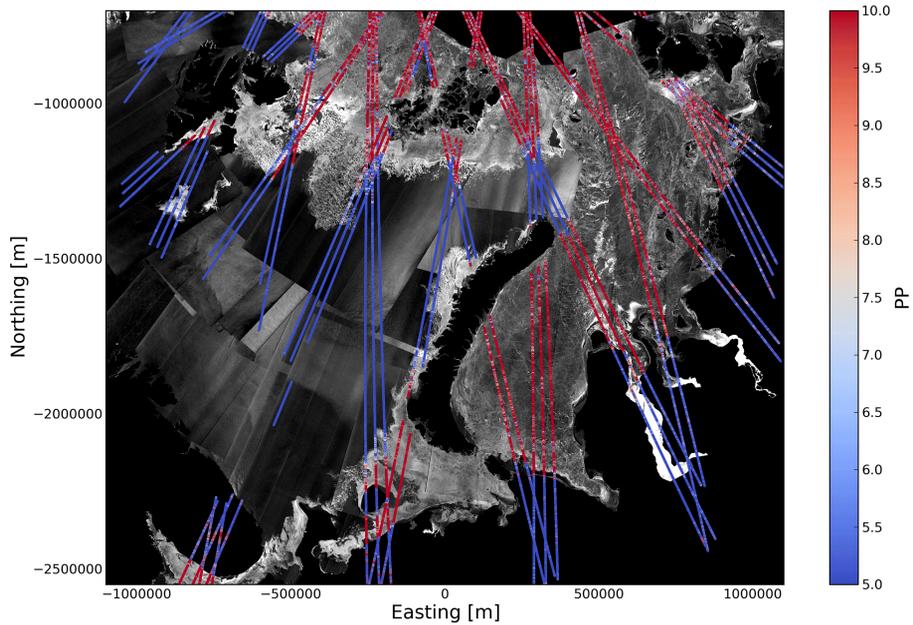


Figure 1. CS-2 Pulse Peakiness (1–5 March 2014) plotted on a RS-2 SAR composite (5 March 2014).

Cryosat-2 operational ice charting

E. Rinne and M. Similä

Title Page

Abstract Introduction

Conclusions References

Tables Figures

◀ ▶

◀ ▶

Back Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



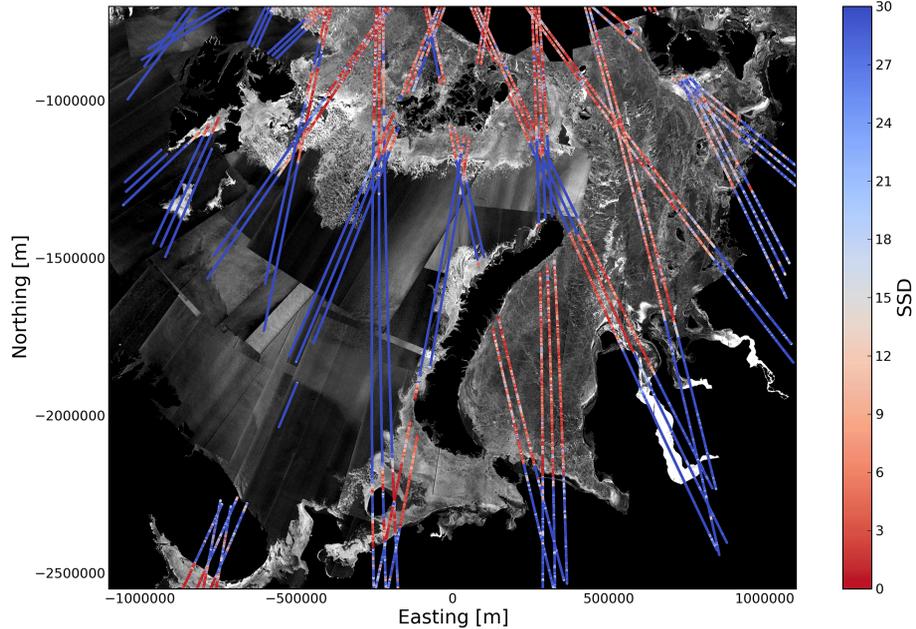


Figure 2. CS-2 Stack Standard Deviation (1–5 March 2014) plotted on a RS-2 SAR composite (5 March 2014).

Cryosat-2 operational ice charting

E. Rinne and M. Similä

Title Page

Abstract Introduction

Conclusions References

Tables Figures

◀ ▶

◀ ▶

Back Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Cryosat-2 operational
ice charting**

E. Rinne and M. Similä

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

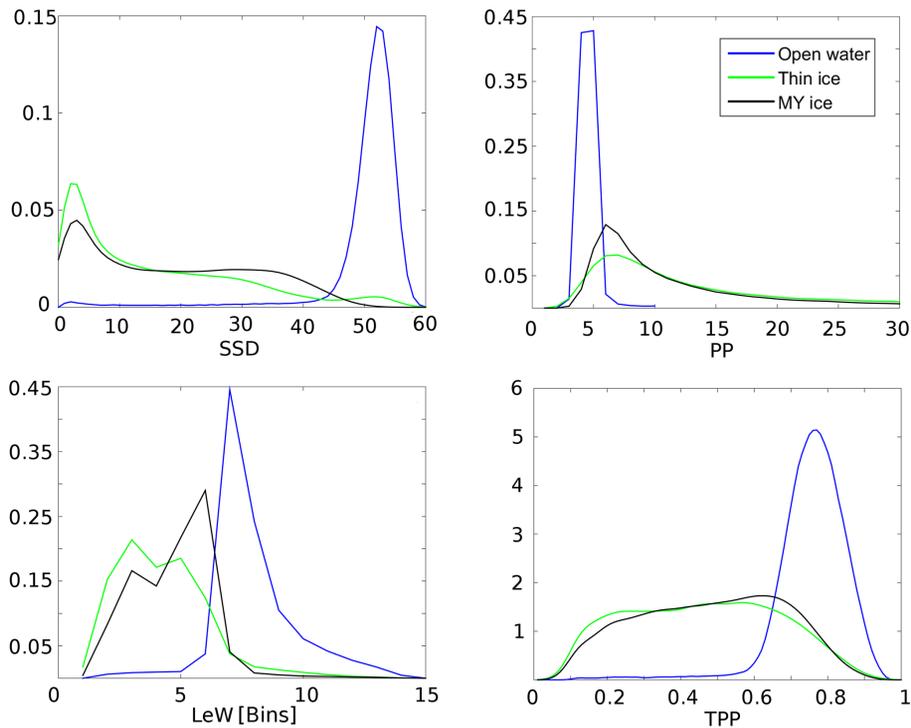


Figure 3. Normed distributions of waveform parameters: SSD (top left), PP (top right), LeW (low left) and TPP (low right) for the period 1–15 November 2013.

**Cryosat-2 operational
ice charting**

E. Rinne and M. Similä

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

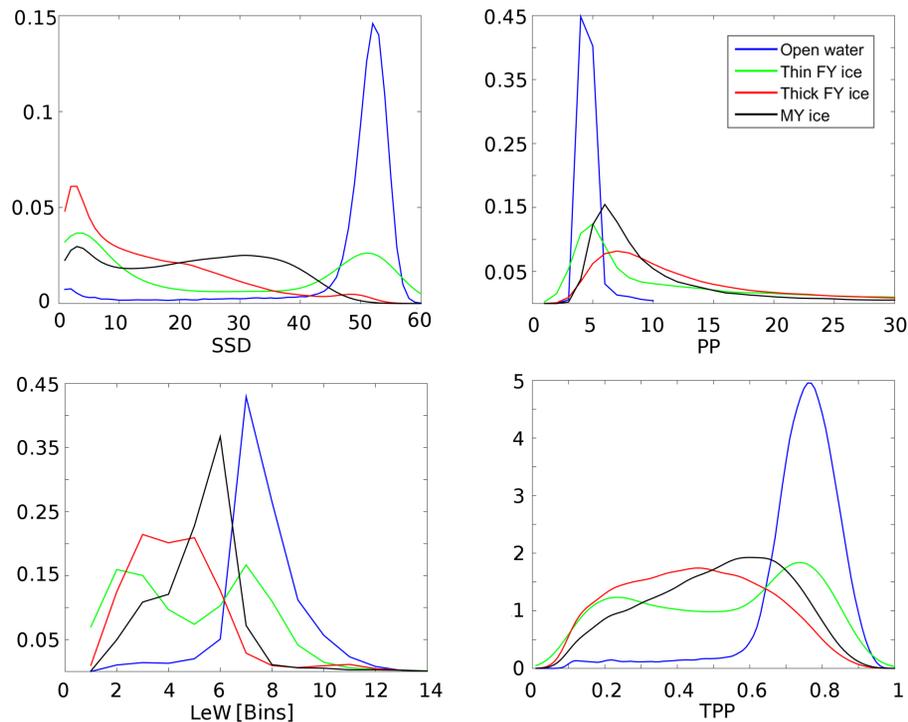


Figure 4. Normed distributions of waveform parameters: SSD (top left), PP (top right), LeW (low left) and TPP (low right) for the period 1–15 March 2014.

Cryosat-2 operational ice charting

E. Rinne and M. Similä

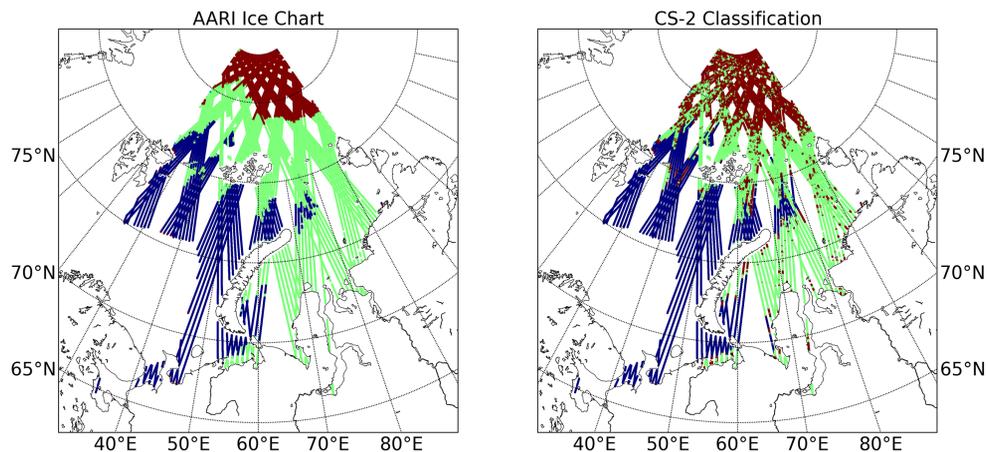


Figure 5. Automatic classification test for March 2014. AARI Ice Chart SA sampled at CS-2 footprints (left) and the classification result from CS-2 measurements (right). 15–30 November 2013. Blue = Open water, green = FY < 70 cm and Red = MY.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Cryosat-2 operational
ice charting**

E. Rinne and M. Similä

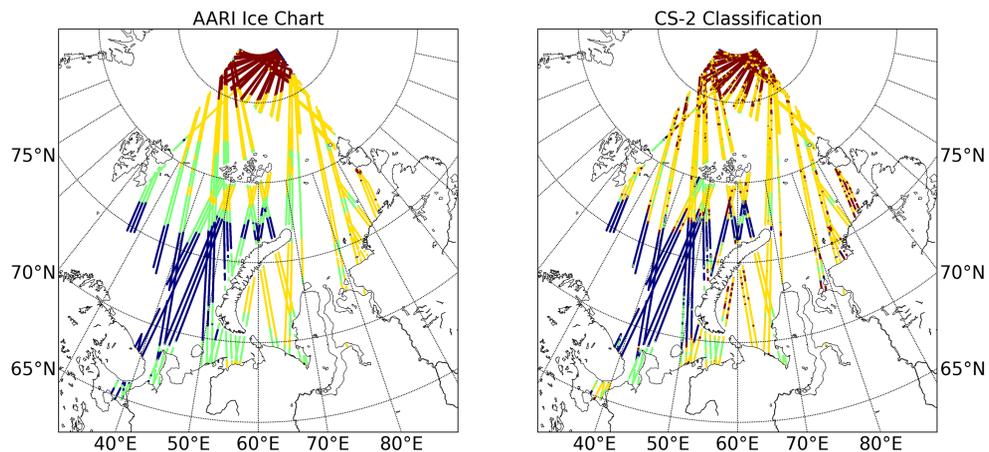


Figure 6. Automatic classification test for March 2014. AARI Ice Chart SA sampled at CS-2 footprints (left) and the classification result from CS-2 measurements (right). 15–30 March 2014. Blue = Open water, green = FY < 70 cm, yellow = FY > 70 cm and red = MY.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)